

Noble Gas Isotope Geochemistry at the Dixie Valley Geothermal Field

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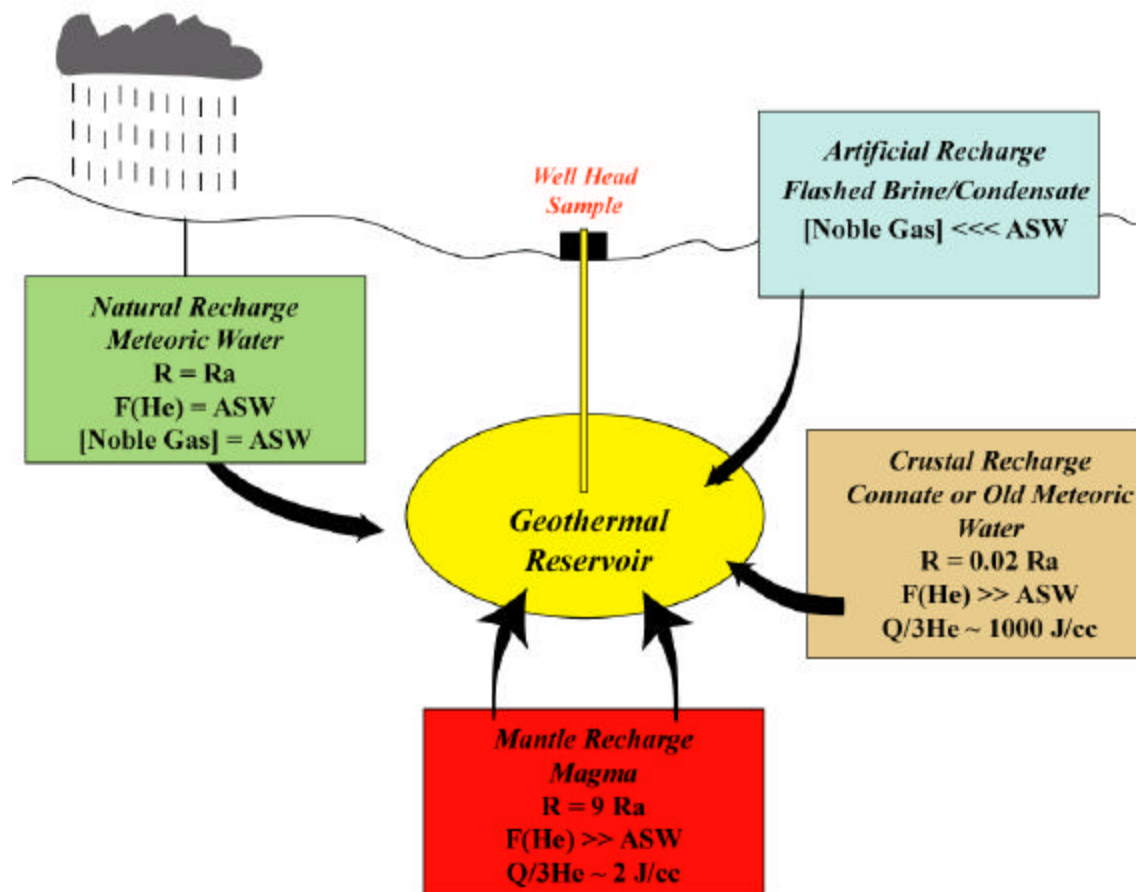
Matthijs van Soest

Primary Goals

- (1) Identify Heat and Fluid Sources**
- (2) Evaluate Noble Gases as Potential Natural Tracers for Monitoring Injectate**
- (3) Integrate Chemical and Isotopic Data into Reservoir Simulation Models**

Noble Gases

Natural Tracers for Geothermal Fluids



**Noble Gases: Sensitive Natural Tracers
For Detecting and Monitoring
Injectate Returns to Geothermal Reservoirs**

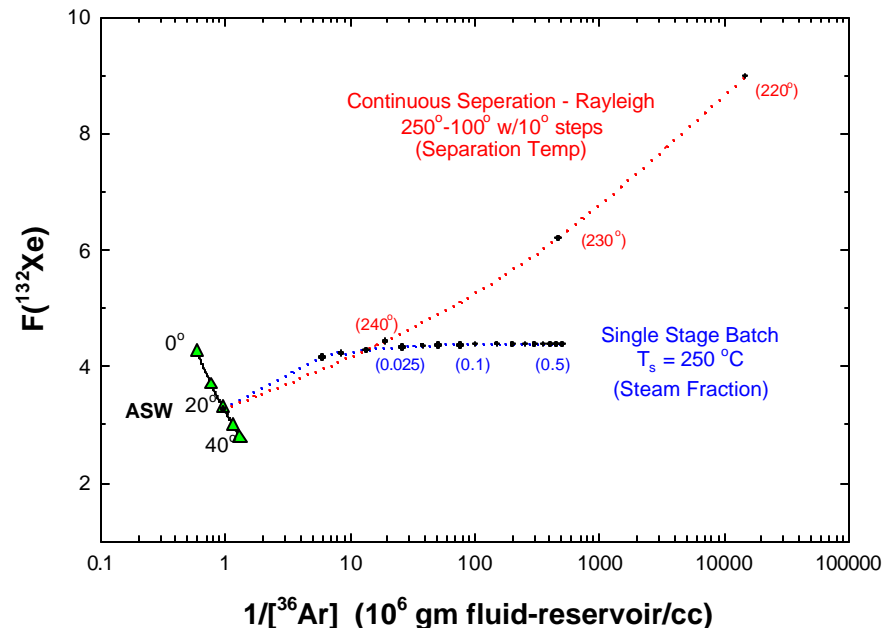
Proof of Concept

- Chloride and Water Isotopes - Widely used
 - Must assume single indigenous reservoir fluid
 - Only applicable in single phase liquid systems
 - Inapplicable in systems with high TDS
 - Low sensitivity: Injectate concentrations are similar to production fluids
 - With 25% steam fraction:
 - $[Cl] \text{ (injectate)} \sim 1.30 [Cl] \text{ (production fluid)}$
 - $D(d^{18}O) \sim 1-2 \text{ ‰}$
- Noble Gases
 - Predictable and relatively invariant composition and concentration in the indigenous reservoir fluids.
 - High sensitivity: Injectate concentrations are extremely low
 - With 25% steam fraction
 - $[Noble \text{ Gas}] \text{ (injectate)} \sim 0.01-0.001 [Noble \text{ Gas}] \text{ (production fluid)}$
 - Noble gases are ~4-40 times more sensitive.

Noble Gases: Tracers for Natural Recharge and Injectate Theory

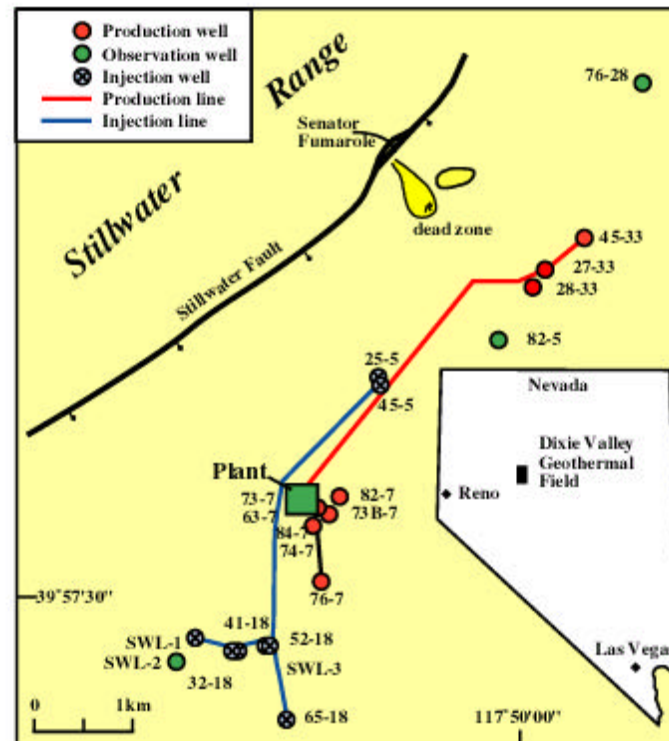


- Phase Separation:
 - ♦ Case I ® Isothermal Batch or Single Stage Separation
 - ♦ Case II ® Non-isothermal Continuous Steam Separation (Rayleigh Distillation).
- Very low solubility leads to high sensitivity for monitoring injectate return.
- With a steam fraction of only 2.5%: residual liquid is depleted in ^{36}Ar by factor of ~20!
- Ultimate composition is path dependent.



Tracers for Re-Injected Fluids at Dixie Valley

Dixie Valley, Nevada Geothermal Field



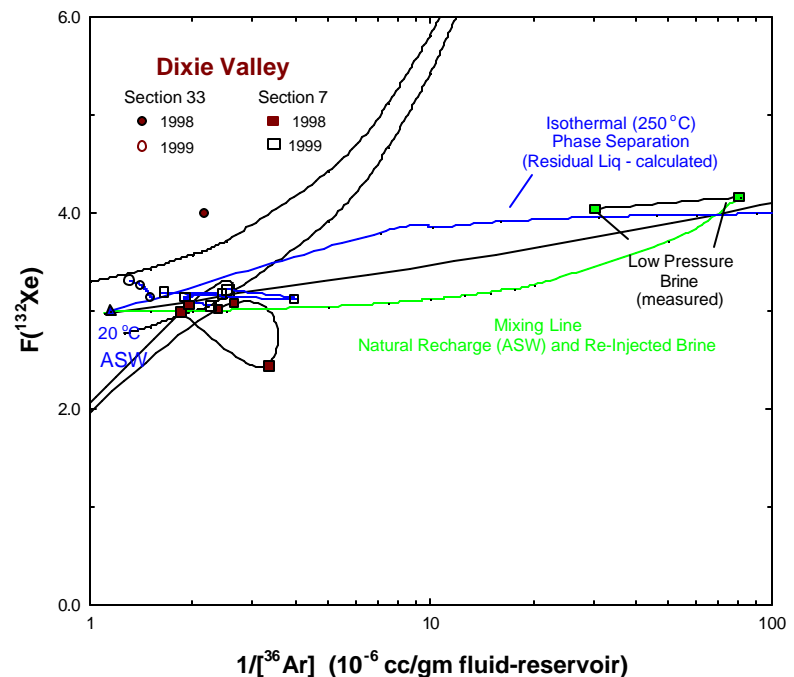
334.5 billion pounds of flashed brine have been injected
into the geothermal field since September 1988

Tracers for Re-Injected Fluids at Dixie Valley

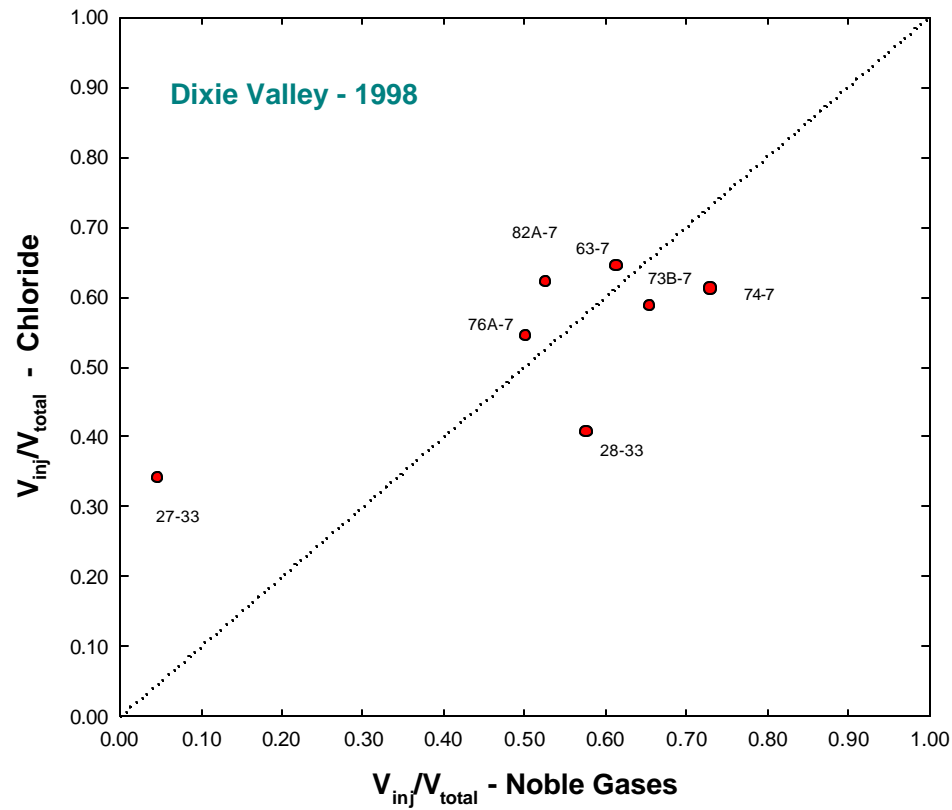


- Composition of re-injected brine is consistent with isothermal batch separation at $\sim 250^\circ\text{C}$ with $\sim 20\text{-}30\%$ steam fraction.
- Noble gases in 1998 and 1999 production fluids are significantly depleted (2-4 times) relative to 25°C ASW.
- Composition of Section 7 wells reflect mixing of re-injected brine and meteoric water.
- Volume fraction of injectate in production stream:
 Section 33 $\sim 30\text{-}35\%$
 Section 7 $\sim 50\text{-}80\%$

Noble Gases: Tracers For Re-Injected Brine



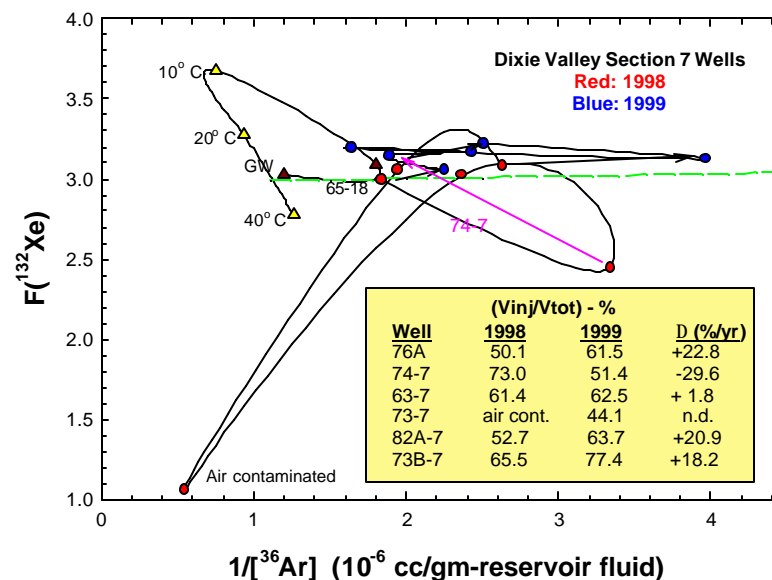
Tracers for Re-Injected Fluids at Dixie Valley



Section 7 Wells 1998 to 1999



- $[^{36}\text{Ar}]$ declined from 1998 to 1999 in all but one well (74-7).
- Relative proportion of co-produced injectate increased at constant rate:
 - $D(\text{Vinj}/V_{\text{tot}}) \sim 20\%/ \text{year}$
- Exception (74-7): $[^{36}\text{Ar}]$ increased by factor of ~ 2 .
 - Cold groundwater added to injectate beginning mid-1997 (Well 65-18)



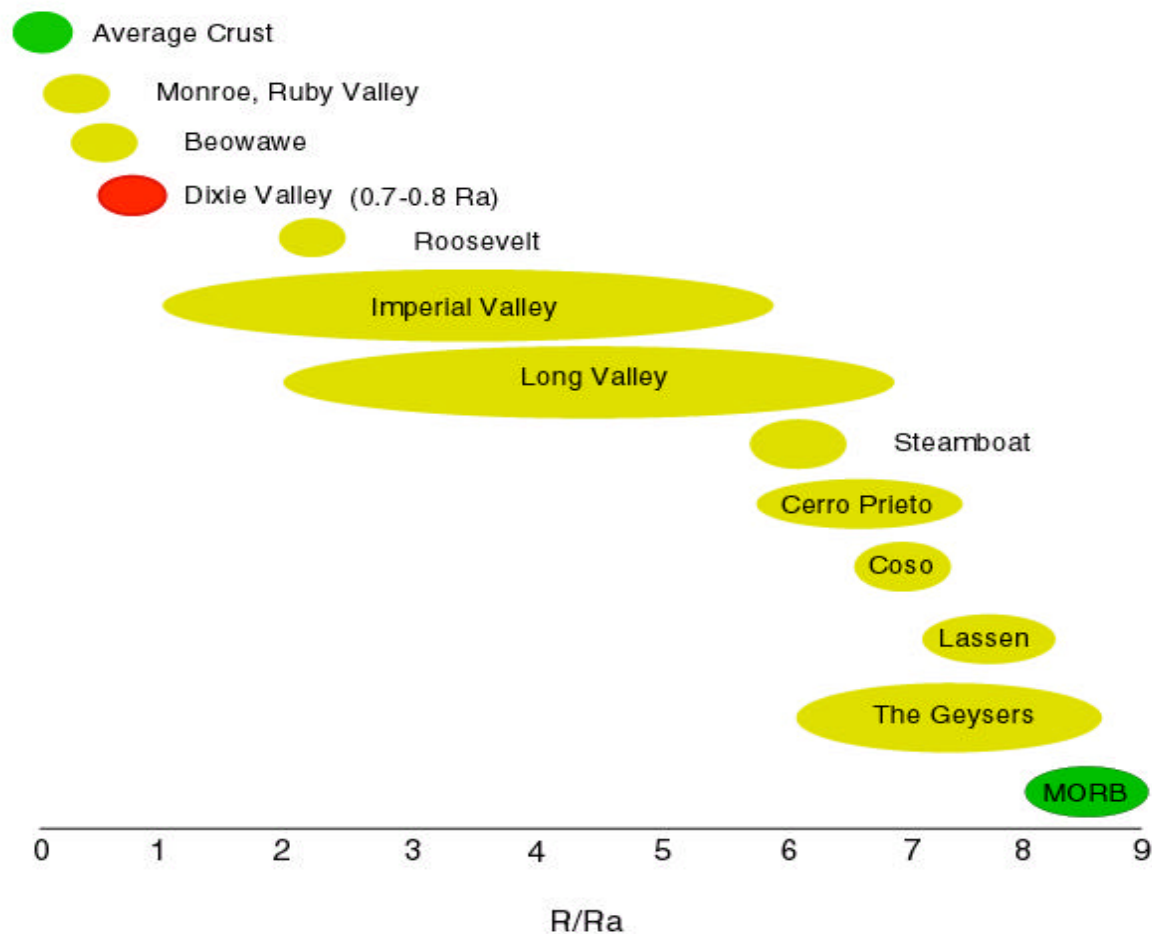
Helium Isotopes in Dixie Valley Wells, Springs and Fumaroles

Heat and Fluid sources

Helium Isotopes in Geothermal Systems



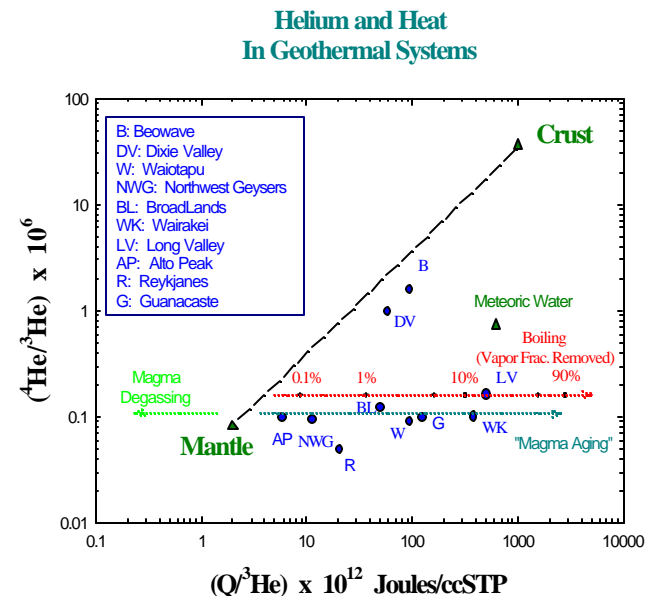
$^3\text{He}/^4\text{He}$ in a Variety of Geothermal Systems



Coupling of Heat and Helium



- ~75% of Earth's heat budget is from natural radio-decay of U and Th --- leads to well defined ($^4\text{He}/^3\text{He}$) and $Q(\text{heat})/^3\text{He}$ ratios for mantle and crustal fluids (green triangles)
- Using this coherence, the heat source of a geothermal reservoir can be evaluated:
 - **Dixie Valley** = 10-15% of heat derived from mantle - remainder is derived from the crustal geothermal gradient
 - **NW Geysers** = 100%
- Heat loss by conduction, boiling, or mixing will shift the helium isotopic composition and Heat/ ^3He ratios in predictable ways --- allowing present state of a geothermal reservoir to be ascertained.



1-D Fluid Flow Model Through Range Front Fault

- Steady state 1-d advection (no dispersion) upward flow scaled to crustal thickness:

$$q = \frac{H_{\text{crust}} * \rho_s * P(\text{He})}{\rho_f * [^4\text{He}]_{\text{f,mantle}}} \left[\frac{(R/Ra)_{\text{meas}} - (R/Ra)_{\text{crust}}}{(R/Ra)_{\text{mant}} - (R/Ra)_{\text{meas}}} \right]$$

q = fluid upflow rate in fault zone

H_{crust} = thickness of brittle + ductile crust

ρ_s, ρ_f = density of solid and fluid

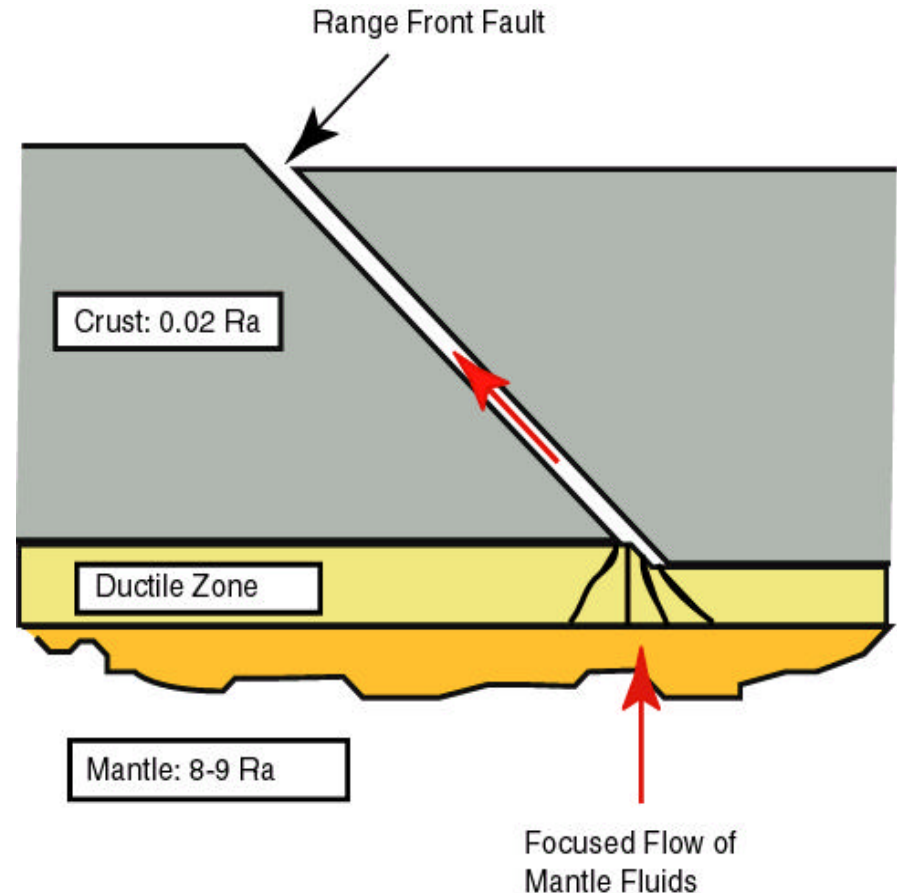
$P(\text{He})$ = present day ^4He production rate from U+Th in fault zone minerals

(R/Ra) = helium isotopic composition

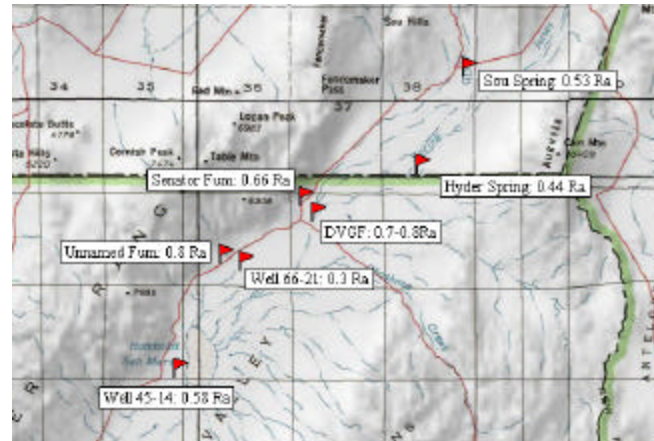
$[^4\text{He}]_{\text{f,mantle}}$ = original ^4He concentration in the upwelling mantle fluid Calculated from ^3He in measured fluid.

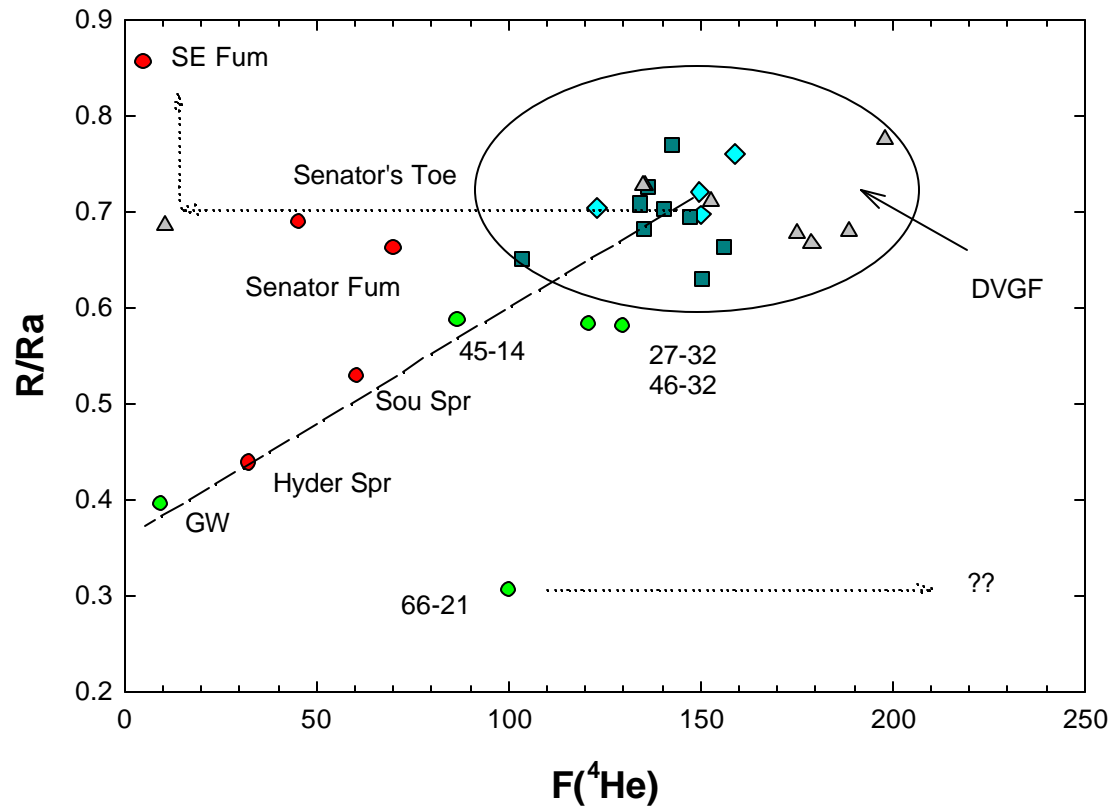
- Dixie Valley geothermal wells ($H_{\text{crust}} = 15$ km; $[U] = 1$ ppm):

$q \sim 0.5$ mm/yr



1-D Fluid Flow Model Through Range Front Fault





Summary

- **Identifying and Monitoring Re-Injected Fluids**
 - ♦ Noble gases compliment traditional conservative tracers by providing a more sensitive quantitative monitoring tool.
 - ♦ Section 7 Wells: ~50-80% injectate and increasing ~20%/year
- **Heat and Fluid Sources**
 - ♦ ~10-15% of heat derived from mantle, remainder from crustal geothermal gradient.
 - ♦ Helium isotopes imply vertical flow rates of mantle fluids through the range front fault of ~0.5 mm/yr.
 - ♦ Helium abundances and isotopic compositions require that Dixie Valley thermal waters are a mixture of shallow young groundwater and a deeper fluid indistinguishable from the fluids produced in the Geothermal field.